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# AD-A205 189

Propagation and Attenuation of Lg Waves in South America

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Observatorio San Calixto Cas 5939 La Paz, Bolivia

31 August 1988

Scientific Report No. 1

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5299

Air Force Geophysics Laboratory

Grant AFOSR-87-0311

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"This technical report has been reviewed and is approved for publication"

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2a. SECURITY CLASSIFICATION AUTHORITY		3. DISTRIBUTION	AVAILABILITY OF	REPORT	
2b. DECLASSIFICATION / DOWNGRADING SCHEDU	LE	Approved fo	or public re	elease; dist	tribution
4. PERFORMING ORGANIZATION REPORT NUMBE	R(S)		ORGANIZATION RE	PORT NUMBER(	5)
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6a. NAME OF PERFORMING ORGANIZATION	6b. OFFICE SYMBOL		ONITORING ORGAN		
Observatorio San Calixto	(If applicable)	All Force	e Geophysics	Laboracory	
6c. ADDRESS (City, State, and ZIP Code)		7b ADDRESS (City	y, State, and ZIP (	Code)	
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La Paz,Bolivia			FB Ma 01731		
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Air Force Office of Scientific Research		Grant AFO			
8c. ADDRESS (City, State, and ZIP Code) Bolling Air Force Base		10. SOURCE OF F	UNDING NUMBER	TASK	WORK UNIT
Washington, DC 20332-6448		ELEMENT NO.	NO	NO	ACCESSION NO
		61101E	7A10	DA	DJ
11 TITLE (Include Security Classification)  Propagation and attenuation of	Lg waves in So	uth America			
12 PERSONAL AUTHOR(S) CABRE Ramón MINAYA Estela	-				
13a. TYPE OF REPORT 13b. TIME C Scientific Report #1 FROM 870	OVERED 1801 TO880731	14 DATE OF REPO		Day) 15 PAGE 30	_
16. SUPPLEMENTARY NOTATION					
17 COSATI CODES	18 SUBJECT TERMS				
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waves, as recorded in LPB star			8.1°W, 3292	m asl, wit	h a gain of
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Epicenter lists were revised, group velocity of Lg phase. So	sismograms insp everal short-pe	sected to ann riod records	were digiti:	ampilitude, zed to draw	particle
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Earthquakes analysed were das	sified according	g to their ch	aracteristi	cs into two	groups
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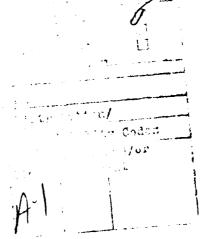
Period of maximal Lg waves changes between 0.5 and 2 seconds. Mean apparent group velocity is 3.6 Km/sec.

Particle motion is characterised by strong polarization horizontal, transverse to the directure of propagation.

For a shield path Lg amplitude is several times that of P; for a cordilleran path Lg is lesser than P, or also it does not appear at all.

# Table of Contents

Introduction. Previous studies	p. 1
Objectives of the present study	3
LPB station	4
Method	4
Data used	5
Analysis	5
Fourier spectrum	10
Interpretation	10
Bibliography	12
Table I Lg velocity in different regions	15
Table II Earthquakes revised	16
Table III Lg mean apparent velocity	21
Fig. 1 Simplified geological map of South America	22
Fig. 2 Map of Lg paths across South America to LPB station	23
Fig. 3 Earthquakes recorded at LPB, travel across the shie	1d 24
Fig. 4 Earthquakes, travel across the shield	25
Fig. 5 Earthquakes, travel across the Andes	26
Fig. 6 Particle motion: a) Brazil earthquake 11-20-80	
b) Colombia 5-14-82	27
Fig. 7 Particle motion: a) Venezuela 7-17-79	
b) Chile 6-28-72	28
Fig. 8 Particle motion: a) Argentina 12-05-77	
b) Peru 6-27-83	29
Fig. 9 Lg spectra, earthquakes from Brazil and Colombia	30
Fig. 10 Lg spectra, Venezuela and Chile	31
Fig. 11 Lo spectra. Argentina and Peru	32



Grant Nº AFOSR-87-0311

PROPAGATION AND ATTENUATION OF LG WAVES IN SOUTH AMERICA

Interim Technical Report

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August 31, 1988

# Introduction. Previous Studies

Lg waves have been studied since 1952.

Press and Ewing (1952) describe for the first time Lg phase, introducing it as a shear wave train of period between 0.5 and 6 seconds, impulsive starting, with maximum amplitude among any other phases. They are recorded after distances of 6000 Km along continental structures. Group velocity amounts  $3.51^{+}_{-}$  0.07 Km/sec, with inverse dispersion, when distance source-station is more than  $20^{\circ}$ . Particle motion is predominantly transverse.

Lehmann (1953) remarked that vertical component of such waves is still significant.

Both (1954) analysed Lg phase, as recorded at Uppsala and Kiruna, noticing also a particle motion transverse. Both remarked two separate windows for group velocity:  $Lg_1$  of  $3.54^{+}_{-}0.06$  Km/sec and  $Lg_2$  of  $3.37^{+}_{-}$  - 0.04 Km/sec.

Ewing et al. (1957), for Hungary and Crimea earthquakes, describe a motion predominantly horizontal transverse, accompanied by a lighter SV motion & coinciding with Both (1954).

Utzu (1960) finds a rather complicated particle motion.

Herrin and Richmond (1960) insist saying that Lg particle motion corresponds to transverse waves trending to some horizontal polariza - tion.

The same again Waldner and Savarensky (1961).

On the other hand Koridalin (1961) shows samples of particle motion with weak horizontal polarization for  $Lg_1$  and strong polarization for  $Lg_2$ .

Pec (1962) focus particle motion as a strongly characteristic of SH.

For epicenters in Iran, Nuttli (1980) finds that the resulting wa ve amplitude is the double of vertical component.

Oliver and Ewing (1957) studied particle motion of  $\rm M_2$ -mode, which is the first mode of shear waves, showing that it should be elliptic retrograde, as actually it was in a quake originated in the Arctic recorded at Palisades.

Group velocities, according to different authors, are listed in table I; it appears that Lg group velocity is confined between a maximum 3.80 Km/sec and a minimum 3.19 Km/sec. Methods of calculating group velocity may change from one author to another and possibly method employed influences the results.

Also frequencies change for different authors (see table I ).

Several investigators find a relation of frequency to distance and to thickness of layer transmitting Lg.

Until now there is no definitive agreement about mechanism both of excitation and propagation of Lg; the main interpretations proposed are the following:

Press and Ewing (1952) introduce Lg as S waves with multiple reflexion within a granitic layer of the continental crust.

For Lehmann (1953) Lg is a superposition of different modes of  $L_{\underline{0}}$  ve waves.

Bath (1954, 1956, 1958) finds energy of  $Lg_1$  generally decreasing with focal depth increasing, energy of  $Lg_2$  is maximum for focal depth about 45 Km, what should be explained through the presence of several guide layers of different velocity.

Gutemberg (1955) suggested the presence of several low velocity guide layers, both in the crust and in the upper mantle.

Herrin and Richmond (1960) state that Lg is transmitted along the

upper crust, one to ten Km (or at most 14 Km).

Oliver and Ewing (1957, 1958), Yamaguchi (1961), Waldner and Savarensky (1961), Brune and Dorman (1963) all consider that a superposition of several higher modes of Love and Rayleigh waves explains frequency contents and group velocity.

Panza el al. (1972) consider that higher Rayleigh modes may be guided both along the crust and along mantle in a layer of a little lower velocity.

Knopoff et al. (1973) demostrated that also Love waves may be transmitted as crustal and mantle channel waves. If no low velocity channel is in the mantle, upper modes of Love and Rayleigh waves should be responsible of Lg waves.

Panza and Calcagnile (1975) arrive to the same conclusion.

Ruzaikin et al. (1977) discard any usefulness of normal modes of surface waves to explain Lg; they suggest that lateral lithological changes along the path influence highly the characteristics of Lg.

Lg appears very relevant in some cases, so that magnitude of a seismic event and some underground characteristics of their path may be revealed; but in other cases Lg waves are not recorded at all; moreover some seismologists are claiming that under the same Lg name we are dealing with different kinds of seismic waves.

Studies of Lg waves across South America are really scarce: Cabré (1969-1971) emphasizes the influence of path on the recording amplitude.

Chinn et al. (1980) focus an especial case of Lg generation converted from Sn at the passage from an oceanic plate to the continental crust.

Raoof and Nuttli (1984-1985) have studied attenuation of Lg in South America.

# Objectives of the present study

The main goals of the present study are:

a) To realize in what instances Lg is significant in the station LPB at La Paz. Bolivia, that is to say, from what seismic foci, across

what regions and with what amplitude Lg phase is clearly recorded.

b) To determine the nature of such waves.

A short synthesis of our results at this point will shed light to the whole report: Lg phase, as recorded in LPB station, is a group of short period waves, significantly large, of frequency 0.5 to 1.6 Hz, relevant for epicenters around the Brazilian or the Guiana Shields, with path mostly along the same shields, with velocity about 3.6 Km/sec, particle motion characteristic of the phase SH.

### LPB Station

LPB is a WWSSN station installed in 1962 (replacing LPZ station 10 cated in a noisy site within La Paz town). Its coordinates are: -  $16^{\circ}31'57.6''S$ ;  $68^{\circ}05'54.1''W$ ; 3292 m asl.

This station has operated quite regularly with a magnification of 50,000 times for short period 3 components until 1979, 25,000 times since then, and 1,500 times always for long period 3 components.

LPB is installed on a thick clay pan.

### Method

During the elapsed first stage of research, the most important work consisted in selecting and observing records to distinguish characteristics of Lg waves, related to group velocity motion, energy spectrum and path from source to receiver.

- 1. Bulletin of the Observatorio San Calixto, Preliminary Determination of Epicenters of the U.S. Geological Survey and Bulletin of the International Seismological Centre were revised to list earthquakes originated in the region deemed to be able to produce Lg at the station LPB.
- 2. Corresponding seismograms were inspected to make a preliminary analysis of Lg phase.
- 3. The time of Lg phase was read and the characteristics of each case annotated, especially amplitude and period.

- 4. Apparent wave velocity epicenter-station was calculated in three ways:
  - a) according to the arrival writen, in some cases, in the bulletin;
  - b) for the wave with maximum amplitude within Lg waves train;
- c) for the first arrival (be it emergent or impulsive), using as a criterion: the higher frequency appearing after longer waves (with some control of long-period seismograms) and the comparison of horizontal and vertical components.
- 5. Digitization of 3-component short-period seismograms for selected events at intervals approximately of 8 points per second, along about 20 seconds.
  - 6. Drawing of particle motion.
  - 7. Fourier analysis.

# Data used

Short period records of 228 South American earthquakes between 1964 and 1986 were considered. Location, origin time and depth were accepted as published by the U.S. Geological Survey as Preliminary Determination of Epicenters (PDE) or in the Bulletin of the International Seismological Centre (ISC).

A few records, 14, were digitized in order to be Fourier analysed and particle motion reconstructed. They were originated:

Brazil 3	Chile 1	
Colombia 3	Argentina	2
Venezuela 3	Peru 2	

### Analysis

Upon a quick inspection of seismograms a pathetic contrast of Lg recording appears for waves transmitted: a) across shields and b) across cordilleran zones (fig. l) So, in the South American map the following zones were considered separately (see table II and fig. 2).

a) Seismic areas with path to La Paz across shield zones:

# a-1. Brazil and Peru-Brazil Border:

14 Earthquakes were considered; the Nos. 1,2,6,9,10,14 have their wave path completely across the Brazilian Shield, N° 3 across the Guiana Massif, Wave arrival to LPB is from East and North-East. They have a Lg emergent but enough clear onset, unequivocally distinguishable from other previous phases.

Lg amplitude increases gradually until reaching a maximum (possi - bly interfering with Rg phase or surface waves, both of long and short period) and then amplitude decreases slowly along a longer duration - (fig. 3).

Maximum amplitude of Lg is three to four times that of P waves. The ratio Lg to S is still much larger, so that often S phase is not detected at all, meanwhile Lg is remarkable.

Lg waves are strongly polarized N-S, so that it is noticed by only an eye comparison of horizontal and vertical records.

Individual characteristics do not show any systematic relation to epicentral distance (10° to 33°), or to focal depth (5 to 65 Km), or to magnitude ( mb 4.7 to 5.5); they do show a relation to azimuth epicenter-station (204° to 250°), without doubt because that yields a different wave path.

Group velocity for the largest waves within Lg was obtained, resulting a velocity higher than that obtained for other regions,  $3.6 \pm 0.2$  Km/sec (Table III).

That high velocity is an effect of the path, not of the origin; it agrees with the nature of rock across a crystalline, ancient basement (also P and S have a higher velocity than in other zones).

# a-2. Colombia and Venezuela

The 94 earthquakes found in Colombia and Venezuela in the interval 1974 to 1984 are of special interest: (Table II) most of their path to LPB station is across the Guiana or Brazilian Shield, (fig, 2) and only a short way beneath the Andes.

Epicentral distance is between 23° and 30°.

Apparent velocity increases with depth (4 to  $160~\mathrm{Km}$ ) as expected, but but

no other depth dependence is found for other characteristics.

The signature of those events is quite different from that appearing in Brazil Zone (fig. 3 and 4). Lg arrival is emergent, with 5 exceptions of impulsive arrival.

Waves are polarized EW.

Wave development may determine two typical envelops; we call them A-or B- type. A-type shows a gradual increment of amplitude until a maxi-mum where the phase Rg may interfere, and after that a slower decrement contrasts with previous increment. B-type does not reach a so large amplitude nor has a relevant maximum.

Generally Lg is larger than P and S (S is so attenuated that often it may not be distinguished on the record).

The mean apparent velocity is for both types 3.6 and 3.58 Km/sec (Ta ble III).

The range of periods extends from 0.9 to 1.6 seconds.

Amplitude ratio Lg/P generally is larger than one (being Lg read at the vertical component and avoiding any confusion with Rg phase, what generally is much larger than Lg).

In some cases Lg is more attenuated than normal, but at the same time P is still more attenuated. In other cases Lg is so strong that it is seen on long-period records; it arrives before Rayleigh waves, several seconds after Love waves (that favors the identification of Lg with higher Love modes).

b) Seismic areas with cordilleran path to La Paz:

### b-l Central Chile

This area is the one with the most problematic Lg in LPB station.

Without doubt it is caused by its path along the Andes, of well known complexity both surficial and subcortical.

No characteristic or general behaviour may be established: Lg phase appears weak in some cases, but in many cases is not apparent at all, not being possible to state that it is not existent or it is so small that it disappears below the S-coda, since it should arrive less than 40 seconds after S and the threshold of readible Lg inscription is greater

than  $m_b = 5$  (in this part of seismogram SS, ScS, etc. are interfering ). However Lg was clearly detected in 10 earthquakes (table 11, fig. 2) with the following characteristics:

Epicentral distance is 10° through 19°; very emergent onset (fig. 5);

Lg may be larger or smaller than P; waves polarized E-W;

Rg is not observed; apparent velocity 3.68 + 0.05 Km/sec (Table III); period between 1.0 and 2.0 sec; some dispersion may be noticed in a few cases;

amplitude or duration of wave train does not increase significantly.

# b-2 Central and South Argentina (Path to LPB parallel to the Andes)

This area is almost as complex as b-1; in the events recorded at LPB station (table II, fig. 2) Lg only hardly may be distinguished; its characteristics are different, but with more regularity, between an event and another.

94 Earthquakes from 1974 to 1984 were analysed, most of them originated in San Juan region (deep foci a priori were excluded, considering null the probability of any Lg phase). Often aftershocks followed the main shock after a few seconds, possibly masking Lg of the main event.

emergent onset; (fig. 5)

Lg is much smaller than P and S waves;

waves polarized E-W;

apparent velocity is 3.52 Km/sec (table III);

short period;

amplitude does not increase or decrease as much as in Brazilian events.

### b-3 Peru and Peru-Ecuador Border

Epicentral distance is 13° through 19°;

Among 14 events occurred between 1982 and 1986 (table II, fig. 2) Lg was recognized in only 7 (including one of epicentral distance less than 10°, considered minimum in other zones). In this case records from the station C.CB sited at about 40 Km from LPB were also considered.

Epicentral distance 8° through 18° emergent onset (fig. 5); amplitude ratio Lg/P generally around one, but not systematic; mean apparent velocity  $3.6 \stackrel{+}{-} 0.3$  Km/sec; period 0.5 through 1.8 sec; amplitude changes are very small; no Rg is observed.

### Generation and Propagation

Lg waves evidently are generated by a constructive interference of SH and SV, arriving to a low velocity layer (possibly the rock of upper most few kilometers); seismic waves are transmitted until recording station. Actual records are a result of both generation and propagation.

Indirect methods are necessary to know in what extent each process contributes to the final recording.

Lg waves period and velocity suggest a thin low velocity granitic layer (possibly 5 km thick), but the absence of that phase in some of the earthquakes of the region rather favors the interpretation that lateral changes of thickness, deep faults (such as Oca fault in N Colombia - Kellog et al. 1982), gradual metamorphism, sedimentary basins are responsible of changing amplitude for seisms of the same region.

# Particle Motion

Lg phase is strongly polarized horizontally for all the paths considered.

So, particle motion is characteristic of all Lg phases, independently of the origin region (figs. 6, 7, 8).

during several seconds of initial Lg waves SH predominates on SV; later on a new arrival, now both at the same time, of SH and SV; after that SV component is stronger than SH.

We shall quote Nuttli (1961) and Nuttli et al. (1962): "According to this relation the surface motion produced by an incident S wave may be considered in two separate cases: 1. Angle of incidence less than  $\sin^{-1}(b_0/a_0)$ , where  $a_0$  and  $b_0$  are the efficient P and S wave velocities at the earth's surface. For this case

there is generated both a reflected P and a reflected S, and the surface particle motion is linear.

2. Angle of incidence larger than  $\sin^{-1}$  ( $b_0/a_0$ ). For this case there is total reflexion of the incident S wave. In general the three components of the earth's surface, namely the SH and the horizontal and vertical components of SV, are out of phase with respect to one another. The surface particle motion is non linear".

# Fourier Spectrum

Spectrum contents (figs. 9, 10, 11) is analysed for all 14 cases as a whole; amplitude logarithm is plotted versus frequency. A maximal pick appears in general between 0.77 and 0.45 Hz; that value does not agree with the first appearence of the record, being the difference originated in the instrumental response curve and in the filtering — used (it amounts about ten percent, and may not influence significan—tly our interpretation in the present research period).

## Interpretation

In general, we have realized for South America what has been said by different authors for other regions:

- 1. Earthquakes orinated well inside an oceanic crust do not produce any Lg phase.
- Cordilleran structure attenuates so much Lg waves that for long path across it Lg either is very small or not existent.
- 3. For surface foci small variations in depth influence strongly Lg generation.
- 4. According to the precedent statement and to the wave velocity, Lg is a train of guided waves in a crustal layer.
- 5. Particle motion maintains some relation both to the angle of  $inc\underline{F}$  dence and P and S velocities

# Specifically we find:

- a) Earthquakes originated North and East of the Brazilian Shield have a prominent phase Lg in the LPB station. Those zones mean politically Northern Ecuador, Colombia, Venezuela, Guianas, Eastern Brazil.
- b) For earthquakes originated in Western Brazil (Brazil-Peru border) Lg is really small, or not existent at all.

- c) Colombia earthquakes have large Lg also for intermediate depth foci; in other zones Lg amplitude decreases rapidly for increasing depth.
- d) Argentinian or Chilean earthquakes, having a path to La Paz mostly beneath the cordillera, do produce really small waves Lg.
- e) Lg appears well transmitted across metamorphic and ingeous rocks not strictly granitic.
- f) Particle motion is characteristic of all phases Lg, independently of the region considered. It is identical to that of SH and well distinguishable since SH arrives before SV and is better developed.
- g) In the case of large Lg waves it seems that they increase by interference of S waves propagation forward and those reflected back in some discontinuity.
- h) For the Venezuela area it is impossible to separate SH and SV; it is not clear if this impossibility arises from the superposition of SH and SV, or possibly from the superposition of Pn, Sn and Love codas.
- i) Being LPB station within a cordilleran zone, though close to its border, the question how Lg waves are recorded so large deserves a longer discussion.

We did not find a convincent explanation of what happens with a train of waves arriving to a medium of high attenuation, but the answer to this question is well known in a similar case: in a local quake the more attenuant is the medium the larger are waves in it (a pathetic example is the town of Mexico destructed by a quake distant about 400 Km, but only destructed in the part where soft underground has a high attenuation). That should mean for our problem that energy conveyed by Lg waves is transformed into shaking amplitude until it is completely dissipated, say not apparent at all after about 200 Km; that contrasts with previos transmision along a regular guide layer, with very little attenuation.

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Table I Lg VELOCITY IN DIFFERENT REGIONS

Table 1 L	Z VELOCITI IN DIFFE	KENT KEGTONS	
Region	Velocity (Km/sec)	Period (sec)	References
Africa S. Africa Transvaal	3.48 to 3.60 3.68 3.66		Gumper y Pomeroy (1970) Willmore et al (1952) Gane et al (1956)
Australia "	$3.50$ $3.44 \pm 0.04$	3 to 6	Bolt (1957) Bolt el al (1958)
Eurasia '	$3.54 \pm 0.06$ Lg <sub>1</sub> $3.37 \pm 0.04$ Lg <sub>2</sub>		Båth (1954)
N America	3.51 ± 0.07	5 to 6	Press Ewing (1952)
Eastern N America	3.57	0.5 to 1 for 19.6° 5 for 35°	Lehmann (1953)
Canadian Shield	3.54		Hodgson (1953)
"	3.60 to 3.70		Brune and Dorman (1963)
11	3.56		Horner et al (1973)
California	$3.54 \pm 0.02$		Press (1956)
Sierra Nevada	3.54 + 0.02 3.53 + 0.02 3.55 - 0.03		11 11
Central Valley and Coastal Ranges	3.55 - 0.03		er er
Central U.S.	3.49		Nuttli (1956)
" "	3.49 to 3.80		McEvilly (1964)
11 11	3.03 to 3.39 2.18 to 3.72		Pomeroy and Novak (1978)
Eastern U.S.	3.19 to 3.35		n n n
" "	3.04 to 3.80		tt tt
Central and SE U.S. SE U.S.	$3.65 \pm 0.04$	0 7 = 0 1	Stauder and Bollinger (1963) Bollinger (1979)
" "	$3.50 \pm 0.13 \\ 3.52 - 0.10$	$0.7 \stackrel{+}{-} 0.1$ $0.8 \stackrel{-}{-} 0.1$	" "
Hungary and Crimea	3.16,3.51,3.23	0.0	Bath (1959)
Coast of Mexico Kamchatka Japan and	3.47 to 3.54	1 to 7	Herrin and Richmond (1961)
Kuril Is.	3.50 Lg		Waldner and Savarenstky (1961)
	3.29 Lg <sub>2</sub>	( ) 10	11 11 11
S China " "	3.53 Lg <sub>1</sub> 3,31 Lg <sub>2</sub>	6 to 10	n n n
Sofia, Bucarest,			
Istanbul "	3.57 Lg <sub>1</sub> 3.32 Lg <sub>2</sub>		Rizikova (1966-68) "
Europe	3.61 3.36		Payo Subiza (1960)
South America	3.8 to 4.4 early I	,g	Chinn et al. (1980)
**	3.6 Lg		f1 11 ff
**	3.5 Late Lg		11 11 11
Black Sea Basin	3.60 Lg 3.55 Lg	15°	Gutemberg (1955) Savarensky et al (1960)
Vicinity of Japan	3.51 Lg <sup>1</sup>		Utzu (1960)
Eurasia and Arctic	$3.54 \text{ Lg}_{1}^{1}$	1 to 9	Koridalin (1960)
Artic	3.54 Lg1 3.48 ± 0.03		Oliver et al (1955)
South West U.S. and	$3.47 \pm 0.01$		Hannin on I Milker (1960)
- Mexico - East coast of Mexico	3.47 - 0.01		Herrin and Milton (1960)
			Herrin and Richmond (1960)
Czechoslovakia	3.2 lg		Pec (1961)

Table II EARTHQUAKES REVISED

No.	Date m d y	Origin t. h m s	Coordinates Lat LongW	Depth m <sub>b</sub>	Δ ^	Az `	Cha	r. Lg/P
Braz	il and P	eru-Brazil						
1	2 13 64	11 21 44.3	18.0s 56.7	16 5.5	10.9	276	e	5.0
	6 13 64	3 56 22.4	2.5N 59.9	65 4.5	20.8	204	е	1.0
	2 22 76	3 24 46.0	0.3N 59.1	10 4.8	19.0	208	e	13.6
	7 13 76	9 21 45.7	7.4S 73.9	33 4.9	10.7	148	e	2.3
	8 24 76	0 14 28.3	8.3S 74.4	90 4.7	10.0	148	e	2.4
	8 - 2 77	17 45 52.5	0.38 50.1	33 4.6	24.2	226	e	4.3
	5 28 78	6 7 3.8	6.7S 74.5	75 4.9	11.6	148	i	3.7
	3 6 80	9 46 17.7	6.1S 71.2	67 4.8	10.8	164	e	10.8
9 1		21 23 5.0	8.18 50.2	33 4.8	19.3	243	е	2.6
10 1		3 29 41.8	4.5S 38.3	0 5.2	31.6	245	i	1.7
	6 21 83	9 23 56.2	8.6S 74.4	152 5.1	10.0	143	e	4.0
	4 8 86	18 2 44.6	7.95 73.9	173 5.8	10.0	147	е	3.7
_	7 14 86	6 34 44.5	9.6S 72.4	33 4.9	8.0	149	е	10.5
14 1	1 40 86	5 19 48.3	5.5S 35.7	5 4.9	33.5	248	е	2.8
Colo	mbia							
15	5 12 75	0 0 39	6.9N 73.1	159 4.6	23.8	168	e	1.3
	6 23 75	5 22 48.3	6.8N 73.1	162 4.9	23.7	152	i	0.6
	2 1 76	3 3 36.3	0.4N 77.2	42 4.6	19.1	152	e	6.0
	3 13 76	21 44 41.8	6.8N 73.0	166 5.3	23.7	168	e	2.0
	5 12 76	16 42 15.1	7.4N 74.9	61 5.1	24.7	164	e	4.4
	6 14 76	1 37 0.1	6.7N 73.0	161 4.8	23.6	168	e	1.2
	8 3 76	2 19 22.7	5.3N 75.9	123 4.7	28.8	160	i	1.1
22	1 10 78	20 8 36.1	3.5N 73.6	42 4.8	20.6	165	e	2.5
	4 28 78	4 28 29.0	12.0N 72.5	13 5.2	28.7	171	e	0.9
	5 15 78	21 52 43.3	6.2N 77.4	6 4.7	24.4	168	i	0.7
25	5 27 78	16 16 42.6	6.8N 73.0	165 4.7	23.6	178	e	1.5
	6 18 78	2 20 25.3	6.8N 72.9	169 4.7	23.7	168	i	0.6
	0 5 78	23 22 21.0	7.4N 76.9	35 4.7	25.3	160	i	0.8
28	5 29 79	12 59 2.5	5.3N 75.7	122 4.9	22.9	161	e	1.1
	9 2 79	2 0 12.4	4.3N 76.4	101 4.7	22.2	159	i	1.7
30	3 6 80	13 42 9.0	6.0N 74.2	60 4.6	23.2	165	e	1.9
31	5 25 80	15 43 30.4	5.4N 74.5	33 5.0	22.5	164	e	1.9
32 1	1 18 80	16 34 38.5	6.8N 72.9	171 4.9	23.7	168	i	1.2
	1 26 80	17 35 41.2	6.9N 72.4	46 4.9	24.6	170	i	3.0
	4 27 81	18 50 38.7	7.0N 76.6	33 4.9	25.0	160	e	1.4
35	5 13 81	4 38 25.0	4.1N 77.1	47 4.6	22.3	157	i	1.8
36	5 20 81	3 1 43.0	4.7N 76.3	136 4.5	22.6	159	i	1.2
	6 13 81	18 39 23.9	6.8N 73.0	171 5.0	23.7	168	е	0.5
	8 5 81	12 58 28.0	3.9N 76.4	62 5.1	21.9	188	e	2.0
	8 30 81	20 50 9.4	6.9N 76.5	35 4.9	24.7	160	е	1.3
	0 24 81	4 24 50.9	6.8N 73.0	167 4.6	23.7	168	e	0.8
	0 26 81	9 5 28.8	6.8N 73.0	165 4.9	23.7	168	i	0.8
42 1		16 15 22	8.85 73.1	62 5.1	9.1	148	i	1.6
43 1	2 17 81	12 54 3.4	6.8N 73.0	165 5.0	23.7	168	6	0.7
44	2 3 82	5 14 36.7	8.15 74.4	169 4.8	10.3	144		0.8
45	2 23 82	20 7 30.8	6.7N 73.0	175 4.7	23.6	168	6	0.8

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Az°Char. Lg/P
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                                8.2N 72.5
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No.	Date	Origin t.	Coordinates	Depth m <sub>b</sub>	Δ°	<b>Az°</b> C	har. Lg/P
	12 25 81	12 35 48.3	10.9N 62.4	96 5.1	27.9	192	e 1.8
94	1 15 82		9.4N 69.9	12 5.1	25.6	176	e 2.0
95	3 18 82		10.5N 62.4	58 4.7	27.4	192	e 1.6
96	5 10 82	1 25 57.3	10.9N 62.5	100 5.2	27.4	192	no
97	5 27 82		8.7N 70.3	14 4.7	25.0	174	e 1.2
98	8 10 82	8 24 0	10.7N 62.6	104 4.8	27.3	192	e 1.3
	11 23 82	17 27 1.0	10.6N 63.2	23 4.8	27.4	190	e 1.2
	12 11 82		8.6N 71.7	14 5.1	25.3	172	e 1.4
101	3 19 83	3 0 26.3	10.6N 63.2	28 4.6	27.4	190	i 1.6
102	4 11 83	8 6 7.2	10.5N 62.7	43 4.7	27.3	191	i 4.0
103	4 11 83		10.4N 62.7	38 5.9	27.3	191	e 1.8
104	5 2 83		10.3N 62.3	45 4.5	27.0	192	e 1.6
105	1 23 84		10.7N 62.7	120 5.4	27.3	191	e 1.7
106	5 25 84	0 59 23.6		41 4.7	27.1	192	i 3.0
107	6 12 84		7.9N 71.3	38 4.7	24.2	173	e 0.9
108	6 14 84	10 4 30.5	9.9N 69.8	38 5.2	26.1	176	e 10.7
Chil	<u>Le</u>						
109	5 15 72	9 12 56.6	29.7S 71.3	49 4.9	13.4	13	e 0.8
110	5 28 72	7 28 13.5	27.7S 71.3	41 4.8	11.5	15	e 0.2
111	5 28 72	9 46 15.4	27.7S 71.4	4 4.9	11.5	16	e 2.3
112	10 26 82	3 24 30.1	29.7S 71.4	63 5.6	13.4	14	e 1.8
113	1 26 86	7 48 23.5	27.0S 70.9	18 5.7	10.8	14	e 3.3
114	3 21 86	13 55 41.0	30.7S 71.4	52 4.9	17.0	13	e 0.7
115	5 14 86	15 54 23.8	32.6S 71.9	33 5.0	17.0	13	e 0.6
116	5 19 86	12 36 30.3	28.45 69.1	113 5.1	15.0	5	e 0.8
117	6 24 86	12 25 28.3	30.75 71.7	51 5.4	17.0	14	e 0.6
118	7 28 86	3 29 56.0	33.3S 72.0	41 4.7	19.0	13	e 0.8
119	7 28 86	20 29 2.7	33.38 71.9	41 5.1	19.0	13	e 0.5
Arge	entina						
120	1 7 74	16 35 56.0	26.9S 65.7	20 5.7	10.5	347	e 0.8
121	8 16 74	7 47 51.5	33.3S 68.3	35 4.8	16.7	359	e 0.8
122	8 24 74	18 58 20.0	31.45 67.4	12 5.3	14.8	357	e 0.8
123	9 3 74	20 22 20.5	25.98 67.6	45 4.8	9.3	357	e 0.8
124		14 40 56.1	30.98 65.1	40 4.7	14.6	348	e 0.8
125		18 10 2.0			16.3		e 0.8
126	1 4 76	4 42 4.0	27.9S 66.0	128 4.8	11.5	350	e 0.8
127	2 14 76	9 1 52	33.6S 68.9	20 4.8	17.1	357	e 0.8
128	3 20 76	2 55 47.8	27.6S 67.4	118 4.8	11.0	356	e 0.8
129	3 27 76	21 5 7.1	31.88 67.7	122 5.1	15.2	358	e 0.8
130	5 4 76	2 7 11.3	27.38 65.8	58 4.7	11.0	348	e 0.8
131	8 3 76	23 43 54.6	31.58 68.5	119 5.0	14.9	358	e 0.8
	10 24 76	0 13 51	32.88 69.2	25 4.9	16.2	356	e 0.8
	11 26 76	7 13 38	28.0S 64.7	25 5.0	11.8	344	e 0.8
134		11 8 43	31.38 67.8	32 5.3	14.7	359	e 0.8
135		11 44 23	31.6S 67.8	60 4.7	15.0	359	Disturbed
	11 23 77	11 46 55.4	31.0S 67.6	8 5.2	14.4	359	e 0.9
	11 23 77	11 58 10.0	31.08 67.9	22 5.5	14.4	159	e 0.9

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No.	Date	Origín t.	Coordintes	Depth	m <sub>b</sub>	Δ°	Az°	Char.	Lg/P
188 1	3 78	1 10 4.4	31.58 67.9	35	5.2	14.9	359	е	0.6
189 1	3 78	6 31 5.1	31.3S 67.9	38	5.0	14.7	359	e	0.6
190 1	17 78	11 33 14.5	31.2S 68.0	20	5.8	14.7	360	e	0.6
191 1	22 78	12 8 26.8	31.58 67.9	35	4.6	14.8	359	e	0.3
192 1	24 78	13 18 15.7	31.7S 68.9	18	5.6	15.1	357	e	0.6
193 3	18 78	0 38 40	31.4S 67.8	16	5.1	14.2	359	e	0.6
194 4	4 78	19 33 53.3	31.2S 67.7	17	5.4	14.6	359	e	0.6
195 5	10 78	23 6 2.0	30.0S 68.9	38	5.1	13.4	357	e	0.8
196 6	7 78	15 16 45.0	32.1S 67.6	44	5.1	15.5	358	e	0.4
197 6	26 78	18 49 11.8	31.6S 67.7	0	5.1	15.0	359	e	0.8
	26 78	1 47 16.1	31.58 67.8	46	5.0	14.9	359	e	0.9
199 8	21 78	0 28 25.1	31.38 67.9	25	5.5	14.7	359	e	0.5
200 10	25 78	1 36 41.7	31.58 67.7	43	5.3	14.9	358	e	0.6
201 1	29 79	3 22 45.5	31.38 68.4	10	5.0	14.7	359	e	0.4
202 7	11 79	23 8 3.1	31.5S 78.3	51	4.8	14.6	359	e	0.6
203 8	30 79	18 59 46.9	31.58 67.7	47	5.4	14.9	358	e	0.8
204 10	8 79	1 52 40.0	31.55 68.0	39	4.8	14.9	360	e	0.8
205 1	17 80	11 0 10.0	31.58 67.7	29	4.8	14.9	359	e	0.8
206 1	24 80	18 34 4.1	31.85 68.5	43	4.7	15.2	359	e	0.8
207 4	9 80	8 17 57.4	31.68 67.5	23	5.4	15.1	358	е	0.7
208 5	25 80	21 46 11.8	31.35 68.0	43	5.0	14.7	360	е	0.7
209 11	10 80	16 24 39.0	31.68 67.5	13	5.6	15.0	358	е	0.7
210 12	6 80	3 43 11.5	31.3S 67.5	46	4.8	14.7	3.8	e	0.7
211 7	2 81	11 3 35	33.0S 69.1	58	4.5	16.4	357	е	0.7
212 8	4 82	5 12 20.2	30.58 68.1	53	4.8	13.9	360	e	0.6
213 12	4 82	3 26 42.6	31.35 67.7	36	4.9	14.7	359	е	0.6
214 5	6 84	13 57 31.7	30.6S 68.9	81	4.6	14.0	357	e	0.6
Perú									
215 8	12 82	8 27 28.0	6.7S 75.8	33	4.7	12.3	143	i	3.1
	15 82	6 11 16.8	10.15 76.5	117	5.5	10.3	129	e	4.0
	27 83	5 5 19.0	13.58 76.8	33	5.4	8.9	111	e	0.6
218 3		1 56 38.6	10.58 74.9	18	5.4	8.9	133	e	1.5
219 3		1 2 51.0	3.75 78.3	46	4.5	16.2	143	e	1.5
220 3		16 47 51.8	8.1S 80.1	33	4.9	15.3	127	no	
221 4	_	6 49 30.4	4.1S 80.8	33	4.7	18.0	135	no	
	23 86	20 27 12.7	3.9S 80.9	46	5.4	18.0	136	e	0.5
223 4		13 43 12.0	15.0S 75.5	57	4.8	8.1	103	e	0.6
224 5		3 56 56.2	3.4S 76.7	121	4.9	18.0	148	e	1.2
225 7		18 11 22.9	9.25 79.9	33	4.9	14.0	124	e	1.6
226 7	20 86	5 52 16.2	15.98 75.3	33	4.6	9.1	96	no	
227 7	21 86	0 40 46.8	14.1S 76.2		4.1	8.0	108		
228 7	27 86	2 54 34.4	10.05 76.3	33	4.7	9.0	130	no	

# Table III Lg MEAN APPARENT VELOCITY

Origen Zone	Velocity (Km/sec)
Brazil	$3.60 \pm 0.2$
Colombia	$3.60 \pm 0.2$
Venezuela	$3.58 \pm 0.2$
Chile	$3.68 \pm 0.3$
Argentina	$3.52 \pm 0.2$
Peru	3.60 ±0.3

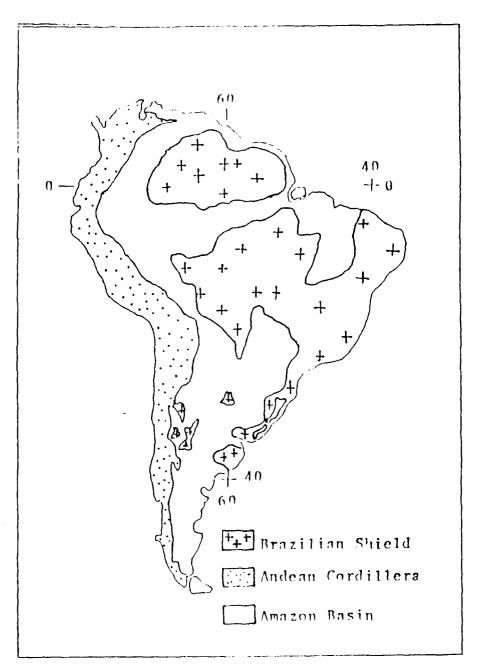


Figure 1: Simplified Geological map of South America. (Martinez, 1980; Almeida and Hasui, 1984)

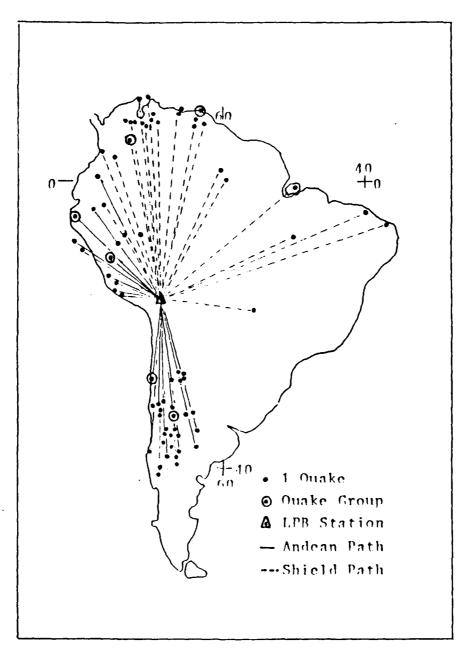


Figure 2: Map of "Lg" paths across South America to LPB station

			man manusco (1777) por monemo, por transfer (Milmore) (Light) (A) (A)	
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			Williams in the second	
			S. The Water	
	e in section of the	1 = 25 K '28/78 nin. 13sec.	And to Like the Complete Comp	
COMP 2 SP A = 50 K BRAZIL 02/22/76 Lg 03h. 34min. 42sec.	Service were the service of the serv	COMP 2 SP A = 25 K BRAZIL 05/28/78 Lg 06h. 13min. 13sec.	(Anti-Mineral/September)	
COMP 3 SP A = 50 K BRAZIL 02/22/76 Lg 03h, 34min, 42sc	Tanks of the state		Yan' Yananya Yanan	

COLOMBIA 03/13/76 Lg 21h 56min 36secThe second of the second of th

Earthquakes recorded at LPB station, travel across the shield. Fig. 3

المراق ال

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	ひじょうさんなくなんないからくるはん	COMP 2 SP A = 25 K VENEZUEIA 06/14/84 Lg 10h, 18min, 02sec.	
COMP Z SP A = 50 K VENEZUELA 12/02/76 Lg 05h, 47min, 52sec.	N. C.		

Earthquakes recorded at LPB station travel across the shield.

COMP 2 SP 4 = 50 K		
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Earthquakes recorded at LPB station, travel across the Andes. Ŋ Fig.

# PARTICLE MOTION

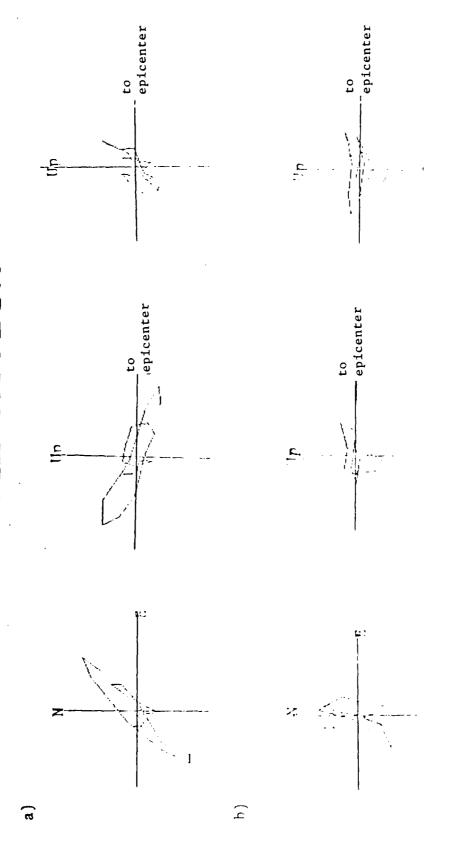


Figure 6: Horizontal and vertical components particle motion diagram:

a) Brazil earthquake

b) Colombia earthquake

11/20/80

5/11/82

# ANTICH HOLLS

3)

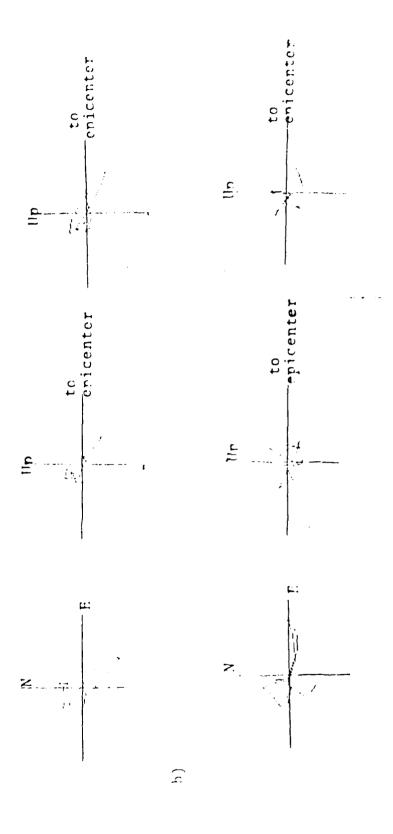


Fig. 7 Horizontal and vertical components particle motiom diagrams:

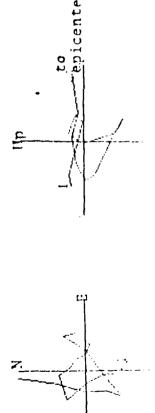
a) Venezuela earthquake 7/17/79

b) Chile earthquake 6/

6/28/72

# PARTICLE MOTION

a)



epicenter

<u>2</u>

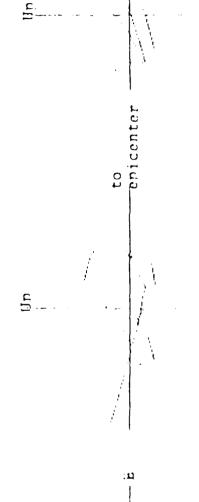


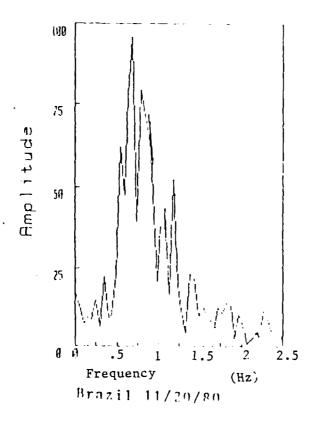
Figure 8: Horizontal and vertical components particle motion diagrams:

12/05/77 a) Argentina earthquake

Peru earthquake

р)

6/27/83



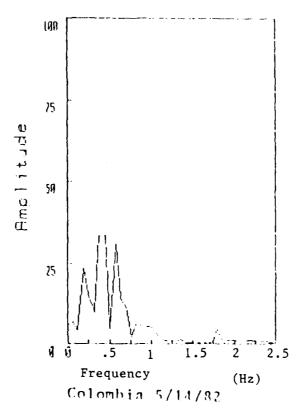
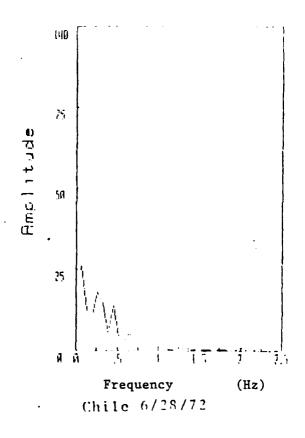


Figure 9: Lg spectra for earthquakes originating in Brazil and Colombia



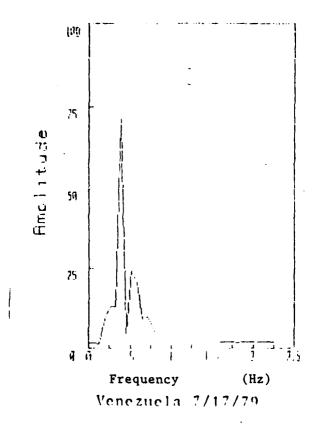
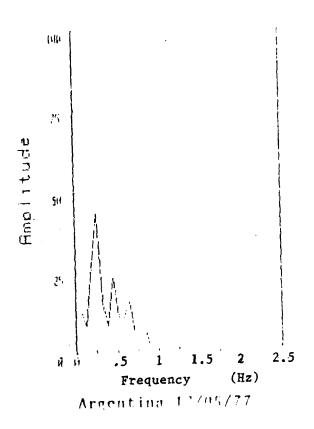
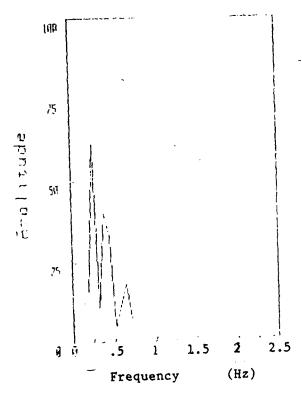


Fig. 10 Lg spectra for earthquakes originated in Venezuela and Chile





Perú 2/27/83

Fig. 11 Lg spectra for earthquakes originated in Argentina and Peru